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# Atmospheric mercury inputs in montane soils increase with elevation: evidence from mercury isotope signatures

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The influence of topography on the biogeochemical cycle of mercury (Hg) has received relatively little attention. Here, we report the measurement of Hg species and their corresponding isotope composition in soil sampled along an elevational gradient transect on Mt. Leigong in subtropical southwestern China. The data are used to explain orography-related effects on the fate and behaviour of Hg species in montane environments. The total- and methyl-Hg concentrations in topsoil samples show a positive correlation with elevation. However, a negative elevation dependence was observed in the mass-dependent fractionation (MDF) and mass-independent fractionation (MIF) signatures of Hg isotopes. Both a MIF ( $\Delta^{199}\text{Hg}$ ) binary mixing approach and the traditional inert element method indicate that the content of Hg derived from the atmosphere distinctly increases with altitude.

**M**ercury (Hg) is a highly toxic metal that can cause serious health problems in humans and animals. It has been widely studied in various environments, such as soils, sediments, and waters. In soils, Hg can exist in different forms, including inorganic Hg (e.g.,  $\text{Hg}^{2+}$ ,  $\text{Hg}^0$ ,  $\text{Hg}^{2+}$ -complexes) and organic Hg (e.g., methylmercury). The bioavailability and toxicity of Hg in soils are influenced by its chemical speciation and isotope composition. Isotope fractionation of Hg in soils has been reported to be affected by various factors, such as soil properties, redox conditions, and biological activity. In this study, we investigated the influence of topography on the biogeochemical cycle of Hg in soils sampled along an elevational gradient transect on Mt. Leigong in subtropical southwestern China. The data were used to explain orography-related effects on the fate and behaviour of Hg species in montane environments. The total- and methyl-Hg concentrations in topsoil samples show a positive correlation with elevation. However, a negative elevation dependence was observed in the mass-dependent fractionation (MDF) and mass-independent fractionation (MIF) signatures of Hg isotopes. Both a MIF ( $\Delta^{199}\text{Hg}$ ) binary mixing approach and the traditional inert element method indicate that the content of Hg derived from the atmosphere distinctly increases with altitude.

K<sub>a</sub> ac ca a a a b c c  
 N<sub>K</sub>, H<sub>v</sub> ( ' a a a c').  
 a a a a a a a P<sub>v</sub> - K<sub>a</sub> H<sub>v</sub> b<sub>c</sub> c  
 c<sup>9,10,13</sup> C , , , c , P<sub>v</sub> b -  
 c ca cc H a c a  
 . I c ca , a b a a a  
 ba / c a H b a  
 a a c - c a a K<sub>a</sub>  
 (., CMAQ-H<sup>14</sup> a GEOS-C<sup>15</sup>), a  
 a a a a a a a K<sub>b</sub> a a c  
 c . I b a a a a a c H  
 a , P<sub>v</sub> K<sub>a</sub> a a a a a a a B a  
 , 27% Ea ' a a c a b c a a ab b 22%  
 P<sub>v</sub> ' a a<sup>16</sup> I C a, a acc P<sub>v</sub>  
 a a a a a a P<sub>v</sub> >50% c a a a a a  
 a a P<sub>v</sub> K<sub>a</sub> >1,000 a.<sup>16</sup> T  
 H a a a a a c C a  
 a P<sub>v</sub> P<sub>v</sub> a a c a a c K<sub>a</sub>  
 a a ba H<sub>v</sub> b a c c .  
 S c a a a K<sub>b</sub> c a b a  
 c c a a a b c a a a  
 a c a a a c c a c P<sub>v</sub>  
 a<sup>8,17</sup> A c a a H a c c  
 14. ac U.S. a P<sub>v</sub> a b  
 H (90%), P<sub>v</sub> b (8%) a ab K<sub>b</sub>  
 b a (<1%). H P<sub>v</sub> K<sub>a</sub> a H c c  
 a a a a c a c a b c a  
 a H c c a a b c a c  
 , a a a a bac , a a c a a a  
 a a c ?  
 O a b a a a H c a b  
 a c , a c a , a c c K  
 a H c c H . T a c a K<sub>b</sub> b a P<sub>v</sub>  
 a a c a c a c H a a  
 c c c a a a c H c c R c  
 a a a H a a K<sub>b</sub> P<sub>v</sub>  
 c a a a a H c a b a c a a c  
 a c c K a a c<sup>19,21</sup> I a  
 a a a a c a (MDF), a a a c a  
 (MIF) a a H<sup>199</sup>H a<sup>201</sup>H a c c  
 a a c c c a a c a a c  
 (MIE)<sup>22,23</sup> a c a K<sub>b</sub> c (NVE)<sup>20,24</sup>. A c b a a  
 MDF a MIF H a a a c a b a a  
 c K a c a c c<sup>25</sup>.  
 M.L (' ') a (2,179 a.,  
 26.39°N, 108.20°E) P<sub>v</sub> Ma Ra G P<sub>v</sub> K<sub>c</sub>,  
 C a(F. S1) a P<sub>v</sub> a c . T a ,  
 P<sub>v</sub> c a a b c a c , , ca P<sub>v</sub>  
 a Na a Na R K (473<sup>2</sup>). B ca , P<sub>v</sub>  
 , P<sub>v</sub> a a K<sub>b</sub> c a a a  
 , P<sub>v</sub> a a c c . T b c M.  
 L a c P<sub>v</sub> a a c c  
 P<sub>v</sub>-S a A , a a b<sup>26</sup> T  
 a K S C a S a (~750 ) ca M.  
 L b a a c b  
 ab a a a (a a c a a ~1250  
 P<sub>v</sub> a >1600 ). F P<sub>v</sub> a  
 a a a , v a a a a c a b  
 0.46°C 100 K a a a P<sub>v</sub> 9.2°C  
<sup>12</sup>. T (~300  
 a ,<sup>-1</sup>, a c <25% a  
 a<sup>11,26</sup>.

## Results

Hg levels and distribution along the elevation gradient. E <sup>V</sup>Ka  
 a H (TH) <sup>V</sup>Ka b K a a c a a  
 P<sub>7</sub> bac aK a C a (0.052 · <sup>-1</sup>(-28)),  
 P<sub>7</sub> a a (a) 0.18 (0.07 0.34) · <sup>-1</sup>a 0.20 (0.08  
 0.38) · <sup>-1</sup>a c c S b 2009 a 2010,  
 c K (F . 1). T a (a) M H c  
 c a 2010 P<sub>7</sub>a 2.16 (0.26 5.05)  $\mu$  · <sup>V</sup>  
 N ca c (ANOVA, > 0.05) TH K  
 P<sub>7</sub>a b K b P<sub>7</sub> a c c a a a  
 2009 a 2010; a a P<sub>7</sub> b c b  
 P<sub>7</sub> c . T TH K ( a 0.19 · <sup>-1</sup>)  
 M . L c a aK ab P<sub>7</sub> b  
 a a U.S. (0.15 · <sup>-1</sup>), N P<sub>7</sub>a  
 (0.19 · <sup>-1</sup>) a S P<sub>7</sub> (0.25 · <sup>-1</sup>)<sup>9</sup>.  
 T TH a M H c c a P<sub>7</sub> a  
 c a P<sub>7</sub> a (<sup>2</sup> = 0.68 0.71, < 0.01 b )  
 (F . 1). T a 0.12  $\mu$  · <sup>-1</sup> <sup>-1</sup>a  
 3.1 · <sup>-1</sup> <sup>-1</sup> TH a M H c K B TH a  
 M H c c a a a c a  
 a a a a a a a a a a a a a  
 F , TH c c a a a a a a a a a  
 b a P<sub>7</sub> K a (<sup>2</sup> = 0.33 0.39, < 0.01  
 b ) (F . S2 S3). T a ca a a a a  
 a ca c ( . <sup>V</sup> a a a c ) a a a a  
 H M . L S K a -a -a aK

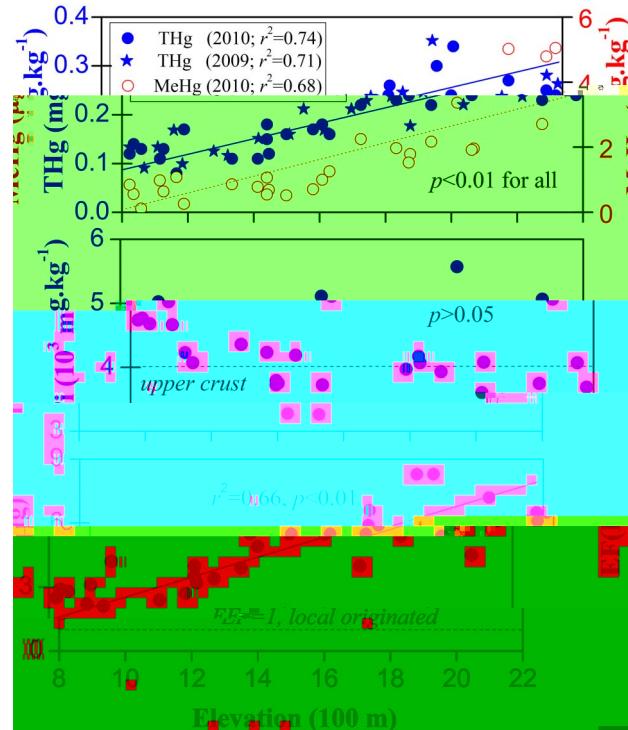
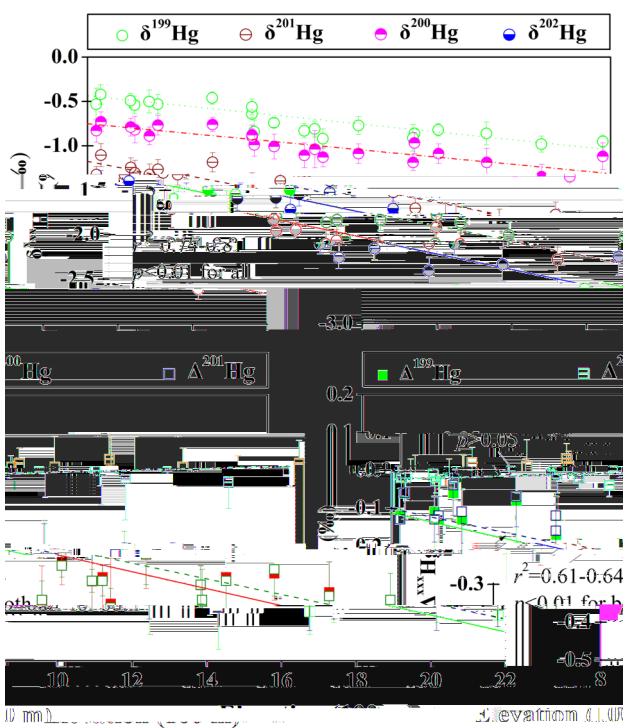


Figure 1 | Scatter plots of soil THg and MeHg contents (upper panel), soil Ti content (centre panel) and calculated enrichment factors ( $EF(Hg) = (Hg/Ti)_{soil}/(Hg/Ti)_{crust}$ ) (lower panel) versus elevation.



**Figure 2 |** Scatter plots of mean  $\delta^{xxHg}$  (upper panel) and mean  $\Delta^{xxHg}$  (MIF, lower panel) isotope ratios in surface soil versus elevation. All error bars represent  $\pm 2$  s.d.

V Ka a c c c a a a H<sup>10,17,18</sup>.  
 H c , a b V c c a a c H a .

**Hg isotope ratios distribution along the elevation gradient.** S. ca. MDF ( $P_f$ ) a 1.2‰ a  $\delta^{202}\text{H}$  ) a MIF (a

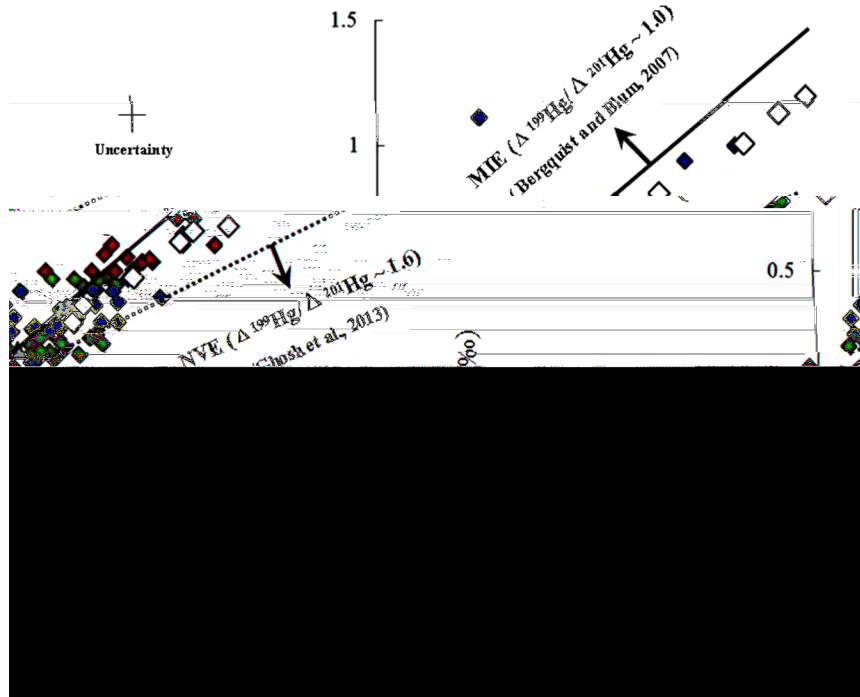


Figure 3 | A comparison of the relationship between  $\Delta^{199}\text{Hg}$  and  $\Delta^{201}\text{Hg}$  from various studies (MIE = magnetic isotope effect; NVE = nuclear volume effect).

0.3% a b  $\Delta^{201}\text{H}$  a  $\Delta^{199}\text{H}$  )  $\frac{v}{a}$   $\frac{p_f}{p_f}$   $b \frac{v}{K}$   
 $\frac{a}{p_f} a \frac{v}{K}$  (F . 2). A c  $\frac{a}{K}$   $\frac{p_f}{p_f}$   $b \frac{v}{K}$   $K$  ( $\chi^2 = 0.61$ , 0.82,  
 $\frac{\chi^2}{K} < 0.01$ ) (F . 2),  $\frac{p_f}{p_f}$  c (%) 100  
 $\frac{K}{a} a a$  b a b a -0.039, -0.040,  
-0.083  $\frac{v}{a}$  -0.083  $\delta \text{H}$  ( $\chi^2 = 199, 200, 201$  a 202,  
 $c \frac{v}{K}$  ) a -0.020 a  $\frac{v}{a}$  -0.021  $\Delta^{199}\text{H}$  a  $\Delta^{201}\text{H}$ ,  
 $c K$  F ,  $\Delta^{201}\text{H}$   $\frac{v}{K}$   $\frac{p_f}{p_f}$   $p_f$  c a  
 $\frac{p_f}{p_f}$   $\Delta^{199}\text{H}$   $\frac{v}{K}$  ( $\chi^2 = 0.98$ ,  $< 0.01$ ) (F . 3). N  $\frac{v}{a}$  ca  
MIF  $\frac{v}{K}$  (.,  $^{200}\text{H}$  a  $^{204}\text{H}$ )  $\frac{p_f}{p_f}$  b  $\frac{v}{K}$  a  
 $\frac{K}{a} a$  (c , a c a ).

## Tracing and quantifying the atmospheric Hg inputs in soil samples. (1). I T Ra

	a	b	$P_7$	H	V	c	M . L
a	T	a	(T), a c	Ka	K		c ca
$P_7$	a	c	$P_7$ a	c	a	c	ca c a
H	c	ac	(EF)	(	a	EF[H])	ba
ac	a	c	c	c	a	<sup>28</sup> acc	
$P_7$	a	:					

$$EF(H) = (H/T) - (H/T) \quad , \quad (1)$$

$$H = H - T \left( H / T \right) , \quad (2)$$

A F . 1 a , EF(H )  $\frac{P}{V}$ a  $\frac{V}{K}$  c a  $\frac{P}{V}$   
 $\frac{V}{K}$ a ( $\chi^2 = 0.66$ ,  $P < 0.01$ ). A EF(H )  $\frac{V}{K}$ a c 1 ca  
 a H  $\frac{V}{K}$ c T c a  
 EF(H )  $\frac{V}{K}$ a  $\frac{P}{V}$ c a  $\frac{V}{K}$ a a - c  
 c H (a  $\frac{V}{K}$ b a H , a c )  
 c a  $\frac{P}{V}$  $\frac{V}{K}$ a T a a c H TH  
 $\frac{V}{K}$ c a a a a 90% a  
 a 50% a ba a ( $\chi^2 = 0.64$ ,  $P < 0.01$ )  
 (F . 4), a ca c b - c  
 c (a a c  $\frac{V}{K}$ ) ac H M.  
 L , a c a a  $\frac{V}{K}$ a .

(2). H a A b B a  
 B<sup>19</sup>, a ca ab H a a c a  
 c a a a H ca b c a ,  
 ac a MDF a ca c a c .  
 C a P MDF, MIF a a c c c a  
 c ca c , a b ab b  
 $\Delta^{199}\text{H} / \Delta^{201}\text{H}$  a . Acc c P , MIF  
 c b a b MDF c , b ca b c a b  
 MIF c v H P  
 MIF a S K a a K b K ca a K  
 $\Delta^{199}\text{H}$  K a c c a a c H  
<sup>29 31</sup>. H P K , a P F . 3, c a K P K  
 c a a a c H c a K P K

$\Delta^{199}$

## Discussion

## Potential mechanisms for Hg isotope signatures in montane soils.

V ca K a H c c c  
 ca b a b ac a H  
 c c c a / H  
 c (., a ca c v ). I a  
 a a c a M.L , K a a (.,  
 a , c a a a a a ) K c  
 b c ca b a K H a a H  
 ac a T ac a H M.L  
 K a a a b a c c -c ca  
 c , c a K a H 0  
 a c H (., c a , a a )  
 a - R H T b R  
 , K a H v a K K c c  
 a - c 22, K a a 24,43 a c b a c  
 H 37.G a , a c c ca c MDF  
 K a H a c H 0 R ca R  
 $\delta^{202}H$  K a a a H 2+. T c H  
 a a MIF H 20,22,44 R a  
 ca MIF c b c K a a a  
 c b a c c 24,37,43 A c b D  
 a 34. a ca a c , K a a a c  
 b a c c b a a c , K a H  
 a a a . I c a a M.L , H  
 b R a c a c  
 ca R a c a 9.S K a a a  
 a a R b c a v a c a  
 c , b a c v , a c a 34.  
 O M.L , ac a K c K H  
 c (., R a ) a a v c c ( a R  
 ). I , H K R c a (0.10  
 $\pm 0.02 \cdot^{-1}, 2 = 2) R R a$

## Potential mechanisms for Hg magnification in montane soils.

W K H , a c a K a a , a c a  
 a a a a a a a b B c  
 a a ' , H a v c c c a c  
 a a c b v K a c . A c a  
 a c a a acc a H a a  
 K a a a a a c F . 5 a a  
 a B c .  
 (1). L a . L a a c c a H a  
 c v a , B c a c a  
 a K c 17. H 0 a v K c

c a K ca R<sub>f</sub> a a a -  
ca R<sub>f</sub> a a a - R<sub>a</sub>  
a a 49,50. T v a c a a -  
H<sup>0</sup> b a K a a a c c K .  
F , H b c c c a  
R<sub>f</sub> ab K a (50% 800%)<sup>17,18</sup>. D ( -  
a ) ca acc 40% a 80% a H v a -  
K c ca <sup>10</sup>, R<sub>f</sub> a H a -  
M.L R<sub>f</sub> a a ba , b a -

- b H a a G (2012)<sup>7</sup> a v ca K H P<sub>7</sub>  
 O M.L , c a c a P<sub>7</sub> a a -  
 c a a a v ( a S I a )<sup>11,12</sup>.  
 S ca K c a b P<sub>7</sub> H K a  
 c a K v a ( <sup>2</sup> = 0.67 0.69, <  
 0.01 b ) (F .S7) P<sub>7</sub> b K v a  
 a a b ca K a c /  
 H - Ka v a a - a c a -  
 a c a c b ab v K.H P<sub>7</sub> K , c a -  
 a a c b a a a c c F  
 a , a a a a b a a a ca ac  
 c H b P<sub>7</sub> a a <sup>59</sup>. S a  
 a a ca P<sub>7</sub> Ka M.L <sup>26</sup>a  
 a ( c a c c K a a c a -  
 b a a )<sup>16</sup>, b c c c a a ,  
 P<sub>7</sub> c , c H a a ac  
 a  
 O M.L , a a a a a a c a a  
 c a , /c a a a c a P<sub>7</sub> c a -  
 Ka <sup>11,12,26</sup>.H c , - ( ' a , c ) , -  
 c a a K c a a a a v c H  
 a c a v H c c a , ca a P<sub>7</sub> Ka . I  
 c a , ca K v a c H b c a a  
 a c a a K c a  
 a a a . T c a b a a a  
 v a c a H c c a a P<sub>7</sub>  
 Ka .

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  - Z a , H, F , X, La , T, S a , L & L, P. B acc a M c K I a c M c R c (O a a K a L) G a . E . 44, 4499 4504 (2010).
  - S a , J, Z , W, L , C-J & F , X. F a ac a v a c c a b P<sub>7</sub> a a ac a a a a K P<sub>7</sub> C . E . 43, 1657 1739 (2013).
  - H , C.D. G ba a c c c c a b b a A . C . P . 10, 12037 12057 (2010).
  - Fa , X, Ob , D, Ha a , A.G, McC bb , I. & Ra , T.H K a ac K a a c b K a a K a ac ab a R c M a a A . C . P . 9, 8049 8060 (2009).
  - H a , J.Y. & G , M.S.EK c a F T . S c M c W D . W U S a . E . 46, 6621 6629 (2012).
  - S a P<sub>7</sub> , C, Ka , J. M. & F a , A.J.T I c a S L a a M c T c D Ac a E K a a G a . E . 46, 8061 8068 (2012).
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**Implications for regional or global Hg cycling.** A a <sup>V</sup> K <sup>V</sup> Ka H  
 a c <sup>V</sup> a b <sup>V</sup> K MDF a MIF a a ac  
 T a ca a MIF ( $\Delta^{199}\text{Hg}$ ) b a a a ac  
 a a a a a a a ca a  
 ac H <sup>V</sup> K a c c a  
<sup>V</sup> a O a a a a a a  
 a <sup>V</sup> K a a <sup>V</sup> K a H ca cc  
 b c a . M <sup>V</sup> c a ca - <sup>V</sup> Ka  
 a <sup>V</sup> K b a a c  
 a , c a , a a ac (., a  
 a a <sup>V</sup> ), O ac <sup>V</sup> a a b c ca O  
 b <sup>V</sup> Ka a <sup>V</sup> K a ba H  
 c c / b a a K ca a a H  
 a a b a , a a acc a  
 ca ba a a a O <sup>V</sup> B  
 a H ab a ca b ac a c H  
 a a <sup>V</sup> T c a b a  
 a a <sup>V</sup> K a ca a a  
 c .

## Methods

TH c c a B<sub>y</sub> a a c a c -  
 Ka. a cab c (CVAAS), a TH c c a  
 a a a a a a B<sub>y</sub> a a a a a a  
 a a a a c Ka a c c c c (CVAFS)  
 c B<sub>y</sub> USEPA 1631.T M H c c a B<sub>y</sub>  
 USEPA 1630.T H c a B<sub>y</sub> B<sub>y</sub> MC-ICP-MS  
 aN-Pa a a c B<sub>y</sub> 12Fa a a c T a  
 c H c c a a a c b a ba, ba  
 , a , c a a a b ca T  
 c b a a a a a a a a a a a a  
 (v ca = 2). W a a a UM-A a a a c a a a (c  
 K 10 a ) a b ac a a NIST 3133. T H c  
 c a UM-A a B<sub>y</sub> a B<sub>y</sub> a a a a  
 ac a a ca  
 Da a a a c , a c c a -  
 a a , a a H c c c a a H  
 a , a , a a a c a c , a c a c a a  
 K S a I a .

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P		A	c a	C	Aca	Sc	c	Na	a Sc c
F	a	C a	(40825011 a	41203092) a	S	-N	B <sub>2</sub>	a c	a
	c	'Ca ac	b	c	c		V	C a - a ca	
G		K c	, B <sub>2</sub>	c P <sub>2</sub> A	b	N	B <sub>2</sub>	a G K	V

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## **Additional information**

**Supplementary information** accompanying the article is available online at <http://www.jbc.org>.

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